## Theoretical calibration of a new paleothermometer based on <sup>13</sup>C-<sup>18</sup>O clumping in carbonate minerals

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Recent measurements of CO<sub>2</sub> molecules containing both <sup>13</sup>C and <sup>18</sup>O (i.e., <sup>13</sup>C<sup>18</sup>O<sup>16</sup>O) at natural, ppm-level abundances<sup>1</sup> have shown that CO<sub>2</sub> formed at low temperatures contains a small, temperature-dependent <sup>13</sup>C<sup>18</sup>O<sup>16</sup>O excess (~1‰ at 298 K). This has made it possible to track sources and sinks of atmospheric gases from a new perspective. Carbonate minerals like calcite (CaCO<sub>3</sub>) are known to retain isotopic signatures over geological timescales, suggesting that measurements of multiply-substituted, or "clumped" <sup>13</sup>C-<sup>18</sup>O-bearing carbonate groups could improve our understanding of ancient and extraterrestrial climates. <sup>13</sup>C-<sup>18</sup>O clumping is an internal property of each phase, so temperature information can be obtained even when the isotopic composition of the fluid phase from which a sample precipitated is not known. This study estimates the abundances of <sup>13</sup>C-<sup>18</sup>O-substituted CO<sub>3</sub><sup>2-</sup> groups in carbonates using *ab initio* thermodynamic modeling, and compares the results to initial measurements of laboratory and natural samples.

Our calculations indicate that carbonate minerals equilibrated at modern earth-surface temperatures will have a slight overabundance of  ${\rm CO_3}^{2-}$  groups containing both  $^{13}{\rm C}$  and  $^{18}{\rm O}$  (i.e.,  $^{13}{\rm C}$   $^{18}{\rm O}^{16}{\rm O_2}^{2-}$ ) relative to what would be expected if carbon and oxygen isotopes were distributed randomly in the crystal lattice. Calcite, dolomite and aragonite will have 0.4%-0.5% excesses of  $^{13}{\rm C}^{18}{\rm O}^{16}{\rm O_2}^{2-}$  at 298 K. Clumping is less pronounced at higher temperatures, disappearing above 1000 K. Experimental results from modern corals and carbonates partially re-equilibrated at higher temperatures suggest an 0.7% excess of  $^{13}{\rm C}^{18}{\rm O}^{16}{\rm O_2}^{2-}$  at  $\sim 300$  K, with a slightly stronger temperature sensitivity than predicted.

<sup>&</sup>lt;sup>1</sup>Eiler et al. 2004, GCA 68:4767.